Overview of Radiation and Arctic Aerosol Interactions with LCLUC

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Where are aerosols in the Arctic?





ARCTIC SWEEP

Cornwallis Island, Nunavut, Canada DAVE BROSHA PHOTOGRAPHY



Local sources of anthropogenic aerosols in the Arctic

- The only significant industrial urban centers within the Arctic are located in Russia.
- \bullet Increasing NO_x concentrations in several Northern Russia cities reflects the increasing number of private vehicles all across Russia.

NO_v concentration in air, µg/m³



SO2 -> Sulfates

Soot (BC)



Local natural and anthropogenic sources of aerosols in the Arctic

• Natural:

- sulfates (gas-to-particle conversion of DMS emitted from sea-water)
- sea-salt (emitted from sea-water)
- organics (emission from vegetation)
- some crustal particulates (from snow-free land surfaces)
- carbonaceous (OC and BC (soot) from wild fires)
- volcanic aerosols
- Anthropogenic sulfates, nitrates and carbonaceous:
 - industrial cities (e.g., metal smelters in Norilsk, Nikel)
 - local use of fossil fuels
 - oil production industries
 - shipping

Local sources are projected to increase in the future





Change in the NOx distribution in the Arctic resulting from ships calculated by the MOZART model (July 2050 - July 2000)



Granier et al. (2007):

- Increases in NO_x from 2000 to 2050: up to 10 ppbv
- Increases in O₃ from 2000 to 2050: up to 30 ppbv
- Increases in Black Carbon from 2000 to 2050: up to 0.1 ppbv (50 ng m⁻³)

Long-range transport of aerosols to the Arctic: "Arctic Haze"



• Winter/spring accumulation of pollution originated at lower latitudes (but to the north of the Arctic Front)

• First reports by pilots in the 50s, measurements from the 70s

- Northern Eurasia and Northern West Europe are major source regions to the BL in the Arctic due to:
 - extension of Arctic front to near 40°N large
 - pollution sources

Understanding of aerosol impacts on the Arctic system requires a knowledge of sources dynamics (LCLUC) in Northern Eurasia, changes in general circulation and climate variability and change.



Modified from Quinn et al. (2008)



Modified from Quinn et al. (2008)

Locations of Long-term Arctic Aerosol Monitoring Stations



Locations of actinometric stations in the Russian Arctic

Operated from the 50-60s until 1993



Measurements of SW and LW total and diffuse radiation

Retrieved aerosol optical depth

Observed long-term aerosols trends in the Northern American Arctic



2005

Observed long-term aerosols trends in Finland

Decrease due to break up of the Soviet Union (Quinn et al.)

Decrease due to introduction of emission control technologies





Time series of SO₂ emissions from the non-metal ferrous² Smelter at Nikel

Yellow bars = Norwegian monitoring station - Svanvik Brown dots = Nikel emissions.

Trends in aerosol optical depth in the Arctic



Trends in aerosol optical depth in the Arctic





Positive values lead to stronger than average westerlies over the middle latitudes

Transport of Asian dust to Alaska







Frequency and intensity of dust outbreaks to the Arctic remain unknown



DOE/ARM North Slope, Alaska

Transport of Asian dust from CALIPSO



Choi, Sokolik, and Winker(2007)

Vertical distribution is a key factor in controlling radiative impacts of aerosols, aerosol-cloudprecipitation interactions and aerosol removal/deposition that affects aerosol transport



Transport of soot to the Arctic





Where are the sources of BC?

Koch and Hansen (2005):

Industrial and biofuel combustion in Southern Asia are a major source of BC in the Arctic

Stohl et al. (2006, 2007):

Western Europe and Northern Eurasia are the main sources of BC



"Non-accounted" summer transport of smoke to the Arctic



Importance of pyro-convective smoke clouds



Aerosol effect on surface albedo



- ice-albedo feedback amplification
- anthropogenic soot may have caused one quarter of last century's observed warming (Hansen et al.)
- significant reductions in Northern hemisphere albedo and sea-ice extent (Jacobson, 2004)
 contribute to melting

Estimates are extremely sensitive to accurate treatment of snowpack aging and soot optical properties as well as modeled predicted soot deposition (Zender et al.)







Effects of surface albedo on radiative forcing of aerosols



mixed vegetation-snow surfaces



Aerosol - clouds interactions



Aerosol chemistry, transport, deposition and aerosol-mixing/ interactions

"Classical view":

Second indirect effect -> precipitation (warm and cold clouds)



Enhanced aerosol amounts can make clouds emit more thermal energy to the surface

(Lubin and Vogelmann, 2006, Nature)



Smaller sizes of drops in polluted clouds



Enhanced aerosol amounts can make clouds emit more thermal energy to the surface (Lubin and Vogelmann, 2006, *Nature*)



Downwelling emission spectra measured by the NSA AERI beneath two clouds with very different condensation nuclei (CN) concentrations.



Trends in downwelling surface LW radiation *Francis and Hunter (2007)*





Dehydration-Greenhouse Feedback (DGF)

Blanchet et al.

•Sulphuric acid coating is observed on aerosol - laboratory observations this indicate coated aerosols inhibit ice nuclei activity by orders of magnitude



- Numerous small ice crystals
- Long cloud lifetime
- Greenhouse warming
- Radar: no or weak return
- Lidar: observed backscatter
- Thin Ice Cloud <u>type 1</u>

- Fewer but larger ice crystals
- Precipitate and dehydrate
- Reduced Greenhouse & cooling
- Radar: visible
- Lidar: Visible
- Thin Ice Cloud <u>type 2</u>





Dehydration-Greenhouse Feedback (DGF)

Clouds forming on acidic ice nuclei precipitate more effectively, dehydrate the air, reduce greenhouse effect and cool the surface



Cold Ice and Snow Surface

Aerosol surface forcings and surface temperature response



Seasonally mean forcing and responses simulated with NASA GISS ModelE GCM

Quinn et al.(2008)

Comparison of Seasonality & Magnitude of Forcing and Surface Temperature Response for Short-lived Pollutants in the Arctic



Quinn et al.(2008)

1 /					
Forcing Agent	Season	Fs	F _{TOA}	F _{TOA-S}	Δ T s ^a
		W m ⁻²	<u>W m⁻²</u>	W m ⁻²	°C
	T	ropos <mark>nheric Aero</mark> sols ·	 Direct Effect[®] 		
Total [*] – Fossil+Bio Fuel	Winter	-0.04	0.08	0.11	-1.4°
$SO_4^{=} + OC + BC$	Spring	0.72	0.92	1.6	-0.93°
	Summer	-0.93	0.11	1.0	-0.47 ^c
	Fall	-0.14	0.08	0.23	-1.1 ^c
	Tr	opospheric Aerosols -	Indirect Effects		
Total [*] – Fossil+Bio Fuel	Winter	$-0.04, 0.24, 0.2^{d}$	0.07, -0.1, -0.03 ^e	0.11, -0.34, -0.23	-0.77 ^f
Cloud albedo + cloud cover	Spring	-3.0, 1.9, -1.1	0, 0.1, 0.1	3.0, -1.8, 1.2	-0.68^{f}
SW, LW, SW+LW	Summer	-12.2, -0.5, -13	6.6, -0.5, 6.1	19, 0, 19	-0.45 ^f
$SO_4^- + OC + BC$	Fall	-0.4, -0.1, -0.5	0.49, -0.9, -0.41	0.89, -0.8, 0.09	-0.89 ^f
Cloud longwave emissivity	Winter	$+3.3$ to 5.2^{g}			1 to 1.6
		Black carbon - Sno	ow Albedo ^h	Soasonal	
BC – Fossil+Bio Fuel	Winter	0.02			0.37
	Spring	0.53		Offset in	0.51
	Summer	0.21		Forcing and	0.21
	Fall	0.002		Beenenee	0.49
	Troposph	e <mark>ric Ozone – GHG wa</mark> l	rmin <mark>g</mark> + SW absorptio	n ⁱ Response	
O ₃ – Fossil+Bio Fuel and	Winter		0.13		0.43
Biomass burning	Spring		0.34		0.31
	Summer		0.14		0.11
	Fall		0.24		0.26
		Methane – GHG	warning ^j		
Methane	Winter		0.29		0.34
	Spring		0.45		0.27
	Summer		0.55		0.15
	Fall		0.34		0.35





- Arctic aerosols may be contributing to the accelerated rates of warming observed in the Arctic. However, there are large uncertainties in assessments because of the inability of models to describe accurately many of the complex processes and feedbacks involved, as well as a paucity of observational data.
- A key question remains as to the role that aerosol, clouds and associated feedbacks play in modulating GHG warming in the Arctic
- By affecting radiation, clouds and surface albedo, aerosols are linked to the radiationclimate feedback processes such as snow/ice-albedo feedbacks, water vapor feedback and cloud-radiation feedbacks that all have been known for some time to be of importance for the Arctic climate.
- Complex spatial trends of aerosol across the Arctic imply the heterogeneous related forcing and complex responses
- A combination of changes in sources (especially, LCLUC in Northern Eurasia), transport (atmospheric circulation) and precipitation (aerosol removal) is emerging as a key factor that controls the presence of aerosols in the Arctic and hence aerosol-induced impacts upon the Arctic system

EXTRA



Global "dimming" paradigm: aerosol-induced reduction in downward surface solar radiation

Trend (%/10YR)

Data of global (direct+diffuse) solar radiation from

160 FSU actinometric stations, 1960 – 1990;

Abakumova et al. (1996)





!!!! Net effect of {Aerosols +clouds +H2O} + surface albedo



Direct radiative impacts of aerosols specifics of the Arctic region				
IMPACT	IMPORTANCE			
Top of the atmosphere (TOA) radiative forcing (SW plus LW)	affects energy balance of the Earth's climate system			
Radiative forcing at the surface (SW plus LW)	affects surface temperature and surface-atmosphere exchange processes, ecosystem functioning			
Radiative heating/cooling (SW plus LW)	affects temperature profile, cloud lifetime, and atmospheric dynamics thermodynamics			
Actinic flux (UV)	affects photolysis rates and photochemistry (e.g., O3)			
Changes in surface albedo (via deposition of soot and dust on surfaces) (SW)	affects surface energy budget, land surfaces and oceans			

Linking aerosol and precipitation through direct and indirect effects





Comparison of measured and simulated surface net shortwave irradiance as a function of visible (500 nm) aerosol optical depth during an Asian dust event at Barrow Observatory, April 2002. Symbols represent one-minute measurements and solid lines the results from MODTRAN[™] fitted using linear regression for zenith angles indicated at the upper right. The circled (suspect) points were not used in the 62° analysis for purposes of computing *DARF* empirically (Stone et al.)





Heating/cooling – Asian dust at Barrow





First Views of CALIPSO data

Browse Image with average every 15 profiles of Total_Attenuated_Backscatter_532

Data Range: 40801: 51600: 1; 1: 583: 1



CALIPSO captures both the maximum altitude of the smoke plume at about 6 km and the vertical smoke structure, as it evolves over time.



AVHRR Arctic <u>summer</u> 20 yr temperature trends (C/year)





Hints at Aerosol-precipitation effects on the climate change of the Arctic

70 year observed trend Hansen, GISS



Specifics of aerosol- radiation interactions in the Arctic



- The deficit of Arctic energy is central to the Globe
- The radiative balance in the Arctic very sensitive to changes of atmospheric composition
- Prolong Polar night importance of LW (aerosols and clouds work like GHGs)
- Because of bright surfaces, even relatively small amounts of absorbing species (such as soot and dust) cause warming
- Changes in

• Important to know: Arctic –wide distribution of aerosol types (composition as a function of size) and their abundance

